

Slip sequences in laboratory experiments resulting from inhomogeneous shear as analogs of earthquakes associated with a fault edge

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ABSTRACT: We explore experimentally and theoretically how fault edges may affect earthquake and slip dynamics, as faults are intrinsically heterogeneous with common occurrences of jogs, edges and steps. In the presented experiments and accompanying theoretical model, shear loads are applied to the edge of one of two flat blocks in frictional contact that form a fault analog. We show that slip occurs via a sequence of rapid rupture events that initiate from the loading edge and are arrested after propagating a finite distance. This event succession extends the slip size, transfers the applied shear across the block, and causes progressively larger changes of the contact area along the contact surface. This sequence of events dynamically forms a hard asperity near the loading edge and largely reduces the contact area beyond. These sequences of rapid events culminate in slow slip events that precede a major, unarrested slip event along the entire contact surface. We show that the 1998 M5.0 Sendai and 1995 Off-Etorofu Earthquake sequences may correspond to this scenario. Our work demonstrates, qualitatively, how a simple deviation from uniform shear loading can significantly affect both earthquake nucleation processes and how fault complexity develops.

1. INTRODUCTION

Faults are often modelled as planar interfaces separating two elastic half-spaces that are driven by a spatially uniform shear that is imposed by the motion of tectonic plates (DAS, 2003; LAPUSTA and RICE, 2003; RICE and BEN-ZION, 1996). Numerous natural faults, however, are loaded by a superposition of the uniform shear due to remote loading (e.g., plate motion), and highly nonuniform loading generated by local fault heterogeneities (e.g. steps, jogs, asperities and edges) (BEN-ZION and SAMMIS, 2003; HARRIS

and DAY, 1993; SHAW and DIETERICH, 2007; WESNOUSKY, 2006) that are abundant along fault-zones. This complexity may govern some earthquake properties.

In this paper we explore experimentally and theoretically the influence of a specific type of nonuniformity which results from the shear loading being applied to a fault's edge. Our laboratory "fault" is formed by two elastic blocks separated by a roughened, but perfectly planar, frictional interface. Shear force is applied to one edge of the slider block (Figure 1) while a uniform normal stress is remotely applied. This loading configuration is an idealized model for the inhomogeneous loading that is likely to occur at an edge or asperity along an otherwise planar fault. To highlight the unique contributions of the nonuniform application of shear to the resulting fault dynamics, this model focuses on the effects of the nonuniform component and ignores the uniform component of the applied shear.

The results of our experimental model are simulated by a spring-slider equivalent that is introduced in Figure 2. A slider of total mass M moves over an immobile rigid substrate (the base). The slider is pushed from its trailing edge (the left side) with a constant velocity V_d through a spring of elastic constant K_d . The force, F , applied to the spring corresponds to the frictional force. It is continuously monitored throughout the simulation. Elasticity of the slider is incorporated by splitting it into N rigid blocks, which are coupled by springs of the elastic constant $K_b=(N-1)K$ so that the slider rigidity is K . This approach is similar to that proposed in the Burridge-Knopoff spring-block model of earthquakes (BURRIDGE and KNOPOFF, 1967), that has been further developed in a number of studies (CARLSON and LANGER, 1989; OLAMI et al., 1992). However, contrary to the most earthquake models where phenomenological laws have been introduced to describe a friction at the slider-base interface, here we explicitly include interactions between each slider's block and the base through an array of "surface contacts", which represent interfacial asperities (BAREL et al., ; BRAUN et al., 2009; BRAUN and RODER, 2002; FILIPPOV et al., 2004b; PERSSON, 1995). As we show below, this model enables us to explain experimental observations and predicts the effect of material properties on the dynamics of the transition to sliding in this type of loading configuration.

Configurations in which edge-loading may play an important role are common along active faults in the earth's crust, and include: slip along a segment within a long fault that loads the neighbouring segments at the edge of the slipped region (e.g. north Anato-

lian fault (STEIN et al., 1997), the physical edges formed between abutting segments (e.g. the intersection of the Susitna Glacier and Denali faults (AGAARD and HEATON, 2004), and by asperities and steps along fault strands (HARRIS and DAY, 1993; JOHNSON et al., 1994 ; LAY et al., 1982; SAGY et al., 2007; SHAW and DIETERICH, 2007; WES-NOUSKY, 2006). As these examples show, the loading of large crustal faults is frequently modelled by a combination of “basal loading” on the crust base, and “edge loading” at the fault edge (LACHENBRUCH and SASS, 1980; MATSUURA and SATO, 1997; RECHES et al., 1994).

2. METHODS

2.1 EXPERIMENTAL METHODS

The experimental setup used is described in detail in (RUBINSTEIN et al., 2004; RUBINSTEIN et al., 2006). We performed real-time measurements of the *true* area of contact, $A(x,y,t)$, along the entire interface separating two polymethyl-methacrylate (PMMA) blocks whose ($x:y:z$) dimensions were $300:30:27\text{mm}$ for the static (“base”) block and either $140:6:75\text{mm}$ or $200:6:75\text{mm}$ for sliding (“slider”) block. x , y , and z are, respectively, the sliding, sample width, and normal loading directions. The initially optically flat base-slider interface was roughened to $1\mu\text{m}$ rms. For the range ($1 < F_N < 4$ kN) of normal load, (F_N), applied, $A(x,y,t)$ varied from 0.35 -1.35 % of the interface’s nominal contact area (DIETERICH and KILGORE, 1994). $A(x,y,t)$ was measured by illuminating the contact area by a sheet of laser light whose incident angle was far below the critical angle for total internal reflection from the PMMA-air layer within the rough interface. Thus light is transmitted only at points of contact, with an intensity at each point (x,y) proportional to $A(x,y,t)$. Using a high-speed camera, we imaged $A(x,y,t)$ at rates up to 100,000 frames/sec. The data acquisition was designed to capture both slow processes at the quasi-static time scales governed by the loading rate and rapid, crack-like, processes whose entire duration takes place in the sub-msec range. As the onset dynamics are governed by one-dimensional rupture fronts (RUBINSTEIN et al., 2004), $A(x,y,t)$ was averaged in y , yielding $A(x,t)$ to 1280 pixel resolution. Thus, $A(x,t)$ provides a local measurement of the contact area, where each pixel measures the integrated contact area a $0.1\text{mm} \times 6\text{mm}$ region (with the higher resolution in the direction of motion). At the initiation of each experiment, before the application of shear the slider was oriented relative to the base to form an initial contact area that was, statistically, spatially uniform (utilizing the $A(x,y)$ measurements to guide the positioning). Upon completion of this initial positioning, F_N was uniformly applied. The corresponding value of $A(x,t=0)$ was then

used to normalize subsequent measurements of $A(x,t)$ to allow us to measure the changes in $A(x,t)$ resulting from the dynamics at each point x .

At $t=0$, a shear force, F_S , was applied in the x direction to one edge (the “trailing” edge at $x=0$) of the slider at a height $z=h$ ($2 < h < 18 \text{ mm}$) above the interface (see Figure 1). F_S was increased from zero at a constant rate (ranging from 1-20 $\mu\text{m/s}$) until, at $F_S=\mu_s F_N$, stick-slip sliding initiated. As long as h was defined as the mean height of the applied shear force the precise means by which F_S was applied was unimportant (e.g. via rigid blocks of various dimensions),. In this loading system (RUBINSTEIN et al., 2007, 2008, 2009), any slip of the trailing edge immediately results in a sharp drop of F_S , which corresponds to the trailing edge of the slider “out-running” the actuator that applied the shear force. While the leading edge (at $x=L$) is stationary, drops of F_S mirror the stress release across the interface.

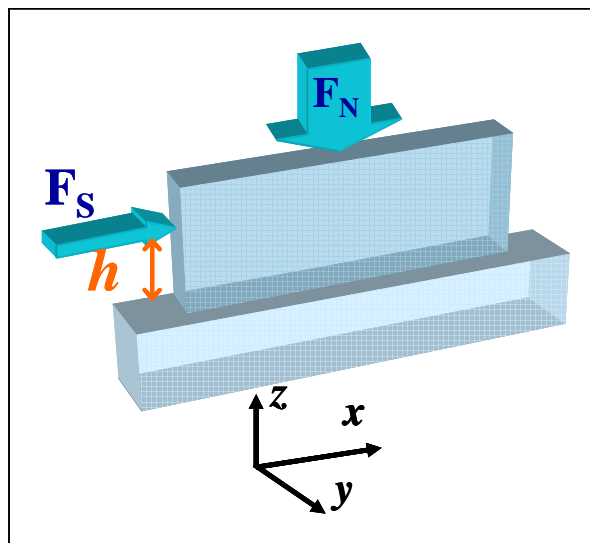


Figure 1: A schematic illustration of the base and slider blocks and load application

2.1 SIMULATION PROCEDURE

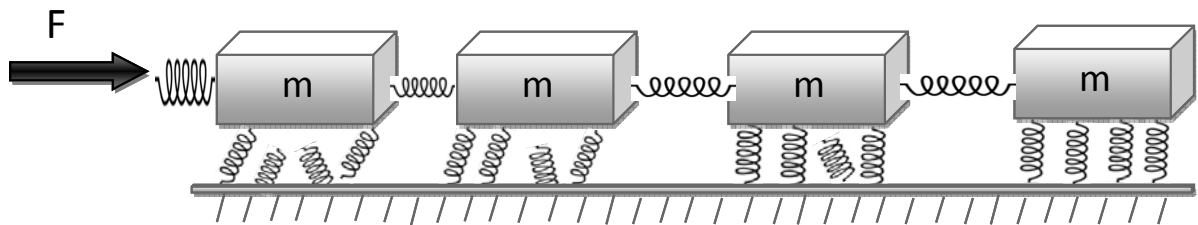


Figure 2: Schematic sketch of a model setup.

Behavior of a single block interacting with a base block through an array of surface contacts been studied previously (BAREL et al., 2010; BRAUN and PEYRARD, 2008; BRAUN and TO-

SATTI, 2009; FILIPPOV et al., 2004a). Here we focus on the collective motion of the chain of elastically interacting blocks. Each contact connects the block and the base through a spring with an elastic constant k_i , where $i = 1, 2, \dots, N_s$ and N_s is the number of contacts between the block and base. The frictional dynamics is governed by two competing processes: (i) formation of contacts (junctions) between the block and the base that tend to inhibit sliding, and (ii) rupture of contacts, i.e., detachment of springs from the block, a process that promotes sliding. So long as a contact is intact, the contact's spring elongates or shortens with the velocity of the corresponding block. Therefore, the interface applies a force on each block of $-f = \sum_i^{N_s} f_i$, where $f_i = k_i l_i$ and $l_i(t)$ is the spring length. Additionally, this force is supplemented by the elastic forces acting from both the left and right neighboring blocks.

To complete the description of the local block dynamics, we must specify the laws that govern the rupture and formation of contacts. We assume that as long as the force, f_i , is below a certain threshold, f_{si} , which models either plastic flow of the contacting asperities or local melting of the boundary lubrication layer, this contact moves together with the corresponding block. When the force exceeds this threshold, the contact detaches from the slider, and reattaches to the block in an unstressed state (i.e., with the length $l_i = 0$) after some delay time, τ . The threshold value, f_{si} , of each contact is assigned a random value from a Gaussian distribution having a mean value $f_s = F_s / (NN_s)$ and a standard deviation Δf_s , where NN_s is a total number of contacts between the slider and base. We note that f_{si} is proportional to the area A_i of a given contact, while the transverse rigidity k_i is proportional to the contact's size, $k_i \propto \sqrt{A_i}$. The distribution of contact's spring constants is, therefore, coupled to the distribution of threshold forces by the relation $k_i = \langle k \rangle (f_{si}/f_s)^{1/2}$, where $\langle k \rangle$ is the mean value of the contact spring constants. When a contact reattaches to the slider, it is assigned new values of the parameters, f_{si} and k_i . The model can be easily generalized to include aging of the contacts by specifying a dependence of f_{si} on the life-time of the contact.

Artificial vibrations of the blocks are avoided by the introduction of a viscous damping force with a coefficient η for the block motion relative to the base, $f_\eta = -m\eta\dot{x}_j$, where x_j is the center of mass coordinate of the j -th block and $m = M / N$ is its mass. We note that the block's oscillations may also be damped due to internal friction within the blocks, i.e., due to phonon excitations inside the slider. The results of our simulations show that the dynamics of transition to sliding are insensitive to any particular choice of the damping force.

We performed calculations over a wide range of the model parameters. In what follows we, however, present results only for a particular set of parameters: $K_d = 4 \times 10^6 \text{ N/m}$, $M = 11.5 \text{ kg}$, $F_s = 1.92 \times 10^3 \text{ N}$, $K = 1.56 \times 10^7 \text{ N/m}$, $V_d = 0.3 \text{ mm/s}$ and $\eta = 0.005 \text{ s}^{-1}$, $D = 0.05$, $\tau = 0.005 \text{ s}$. The values of the macroscopic parameters, K_d , M , F_s , K and V_d were chosen to model the experimental conditions. As for the interface rigidity, $K_s = N_s \langle k \rangle$, we present results for two cases: (i) a soft interface with $K_s = K$ (Figure 7) and (ii) a stiff interface with $K_s = 50 K$ (Figure 10). The interface stiffness and mean rupture threshold force are expected to be directly related to the applied normal load. Most simulations were performed for seventy blocks ($N=70$) with 100 contacts between each block and the base ($N_s=100$).

3. RESULTS

3.1 Slip sequences and stress history

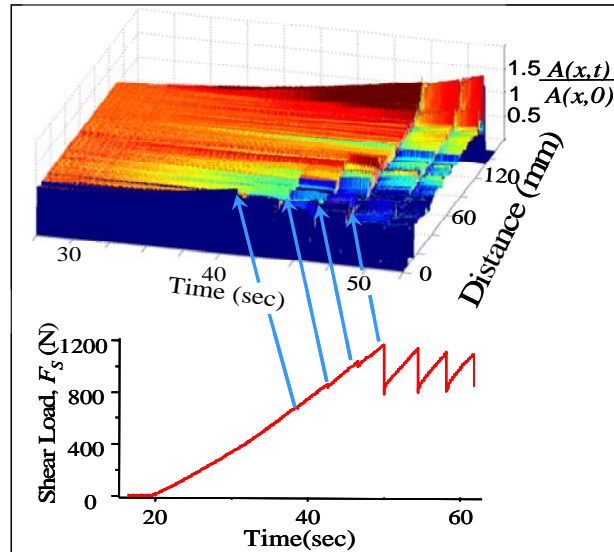


Figure 3: The transition to stick-slip sliding is preceded by a sequence of discrete arrested slip events. (bottom) Applied shear force, F_s , as a function of time with the (top) corresponding spatio-temporal evolution of $A(x,t)$. Each small discrete drop in F_s corresponds to a rapid slip event that arrests within the interface. Each arrested slip event generates significant changes in $A(x,t)$. Arrested events, together with the corresponding drop in shear load, are denoted by the arrows. The large stress drops mark stick-slip motion. Interface length, $L=140 \text{ mm}$, $F_N=3.3 \text{ kN}$.

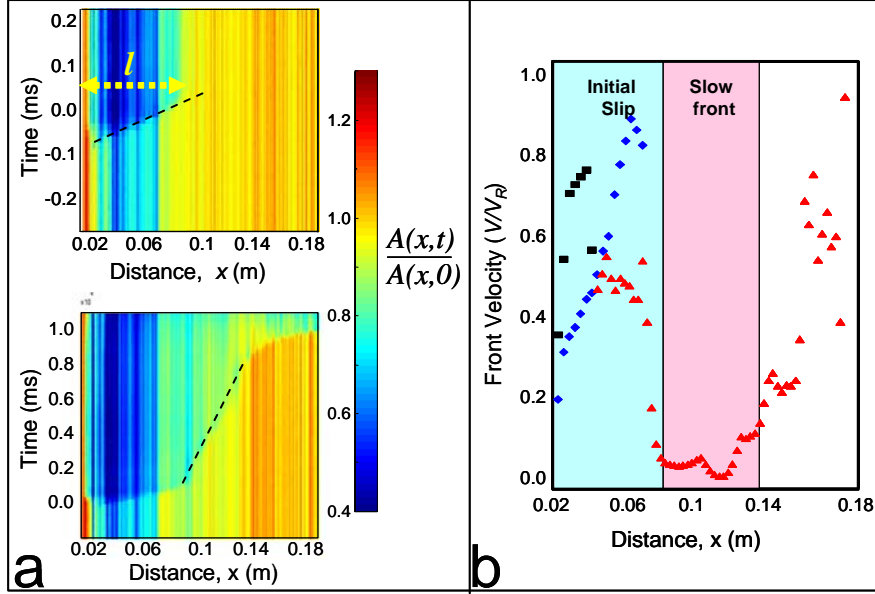


Figure 4: **a**, Time-distance maps of the normalized contact area (scale bar) during slip events across the 0.18 m long experimental fault. (top) An arrested slip event of length, l , generated at stresses well below the onset of stick-slip. (bottom) The transition to stick-slip motion at the peak value of F_S . Here rapid slip arrests and triggers a slow front. At $x=0.14\text{m}$ the slow front nucleates a rapid slip that traverses the remainder of the interface. Color bar indicates the change in $A(x,t)$, relative to the initial, uniform value $A(x,t=0)$ when $F_S=0$. $A(x,t)$, measured at $14\mu\text{sec}$ intervals in two different events. **b**, The slip propagation velocities (as a function of x) of arrested events (diamonds and squares) and the transition (triangles) to stick-slip motion. Velocities are determined from the slope of the event fronts (dashed lines in **a**). Events depicted by diamonds and triangles correspond to (a).

Concurrent measurements of contact area $A(x,t)$, and shear load $F_S(t)$ for a typical experiment (Figure 3) reveal that large, system sized, stick-slip events are the culmination of a complex history of precursory slip events. The $F_S(t)$ curve in Figure 3 reveals a discrete sequence of small sharp stress drops that occur at stress levels well below the peak values of $F_S(t)$. These small stress drops (of $\sim 0.01-0.02 \cdot F_S$) result from the propagation of a sequence of rapid, crack-like arrested slip events. In the short-time measurements of $A(x,t)$ (Figure 4a) we observe that the initial slip events always initiate at the system's trailing edge, and propagate at “sub-Rayleigh” speeds, typically between 60%-80% of the Rayleigh wave speed (V_R), before abruptly arresting (Figure 4). These initial events (Figure 3 – top) are associated with slipping segments of length l that are relatively small compared to the entire fault size, L . We find that l obeys a linear scaling relation (RUBINSTEIN et al., 2007), $l \propto F_S L/F_N$ (Figure 5a). Once l approaches $0.4-0.5L$, this scaling breaks down and the initial dynamics undergo a qualitative change that denotes a transition (RUBINSTEIN et al., 2004; RUBINSTEIN et al., 2006) to a new stage of dynamics. Here, the slip events also initiate at the trailing edge as rapid sub-Rayleigh slip events, but do not simply arrest. Instead, these larger events trigger a “slow” front that

propagates over a range of speeds ($\sim 50m/sec$ in Figure 4b) that can be slower than the sub-Rayleigh velocities of the triggering events by more than an order of magnitude. These slow fronts can propagate stably for some time, and either traverse the remainder of the interface or transition back to sub-Rayleigh fronts, as shown in Figure 4a (bottom). Significantly, overall motion (sliding) between the blocks initiates only after either a slow or subsequently triggered sub-Rayleigh front has reached the leading edge.

The discrete sequence of such arrested slip events, described by Figure 3, is *only* observed when shear is imposed at the sample's trailing edge. It is not observed when, for example, a uniform shear stress is imposed at a remote boundary parallel to the interface. With the trailing edge loading, each sequence initiates via a slip event of finite length, l_0 , with l increasing by discrete increments, Δl , of constant length for each successive slip.

Figure 5b shows that the size of both l_0 and Δl is proportional to h , the height above the interface where F_S is applied at the edge. Note that h , however, has no effect on the overall scaling of l (Figure 5a). Since l scales linearly with F_S , the fixed value of Δl (for a given h) indicates that slip events occur at fixed intervals, ΔF_S , in F_S . Thus, h , which determines Δl , governs also the magnitude of the force drop intervals, ΔF_S , between successive events. For a constant shear loading rate, (as in our experiments) the temporal periods between events are proportional to h .

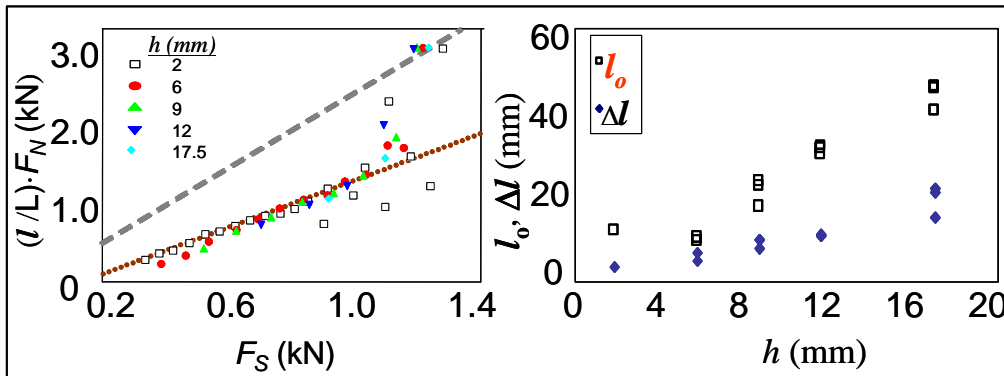


Figure 5: Scaling behaviour of the lengths l of successive slip events. **a**, $l/L \cdot F_N$ as a function of F_S , where F_S was applied at different heights, h , above the interface. h does not influence the $l \propto L \cdot F_S / F_N$ scaling (RUBINSTEIN et al., 2007) (dotted line). Here, $F_N = 3kN$ and $L=140mm$. This scaling breaks down at the transition to large events leading to stick-slip motion, described by the Amontons-Coulomb law (dashed line). **b**, The values of both the initial slip length, l_0 (squares), and the incremental extension of each slip event, Δl (diamonds), increase linearly with h . l_0 saturates at low h suggesting that a minimal length is needed for development of instability. Different points at the same h correspond to different F_N .

3.2 Contact area and fault strength

We now consider the evolution of the contact area $A(x,t)$. Prior to the first event, $A(x,t)$ is spatially uniform. The passage of each successive precursory slip event (Figure 6) significantly al-

ters the contact area, and hence changes the local fault strength. With each successive event, the contact area increases in a region of width D , that is adjacent to the sample's trailing edge. Thus, this process *dynamically* forms an asperity (a localized area whose resistance to slip is much greater than its surroundings (LAY et al., 1982)). Although this asperity continuously strengthens with F_S (Figure 6b), its size, D , remains nearly constant throughout each experiment.

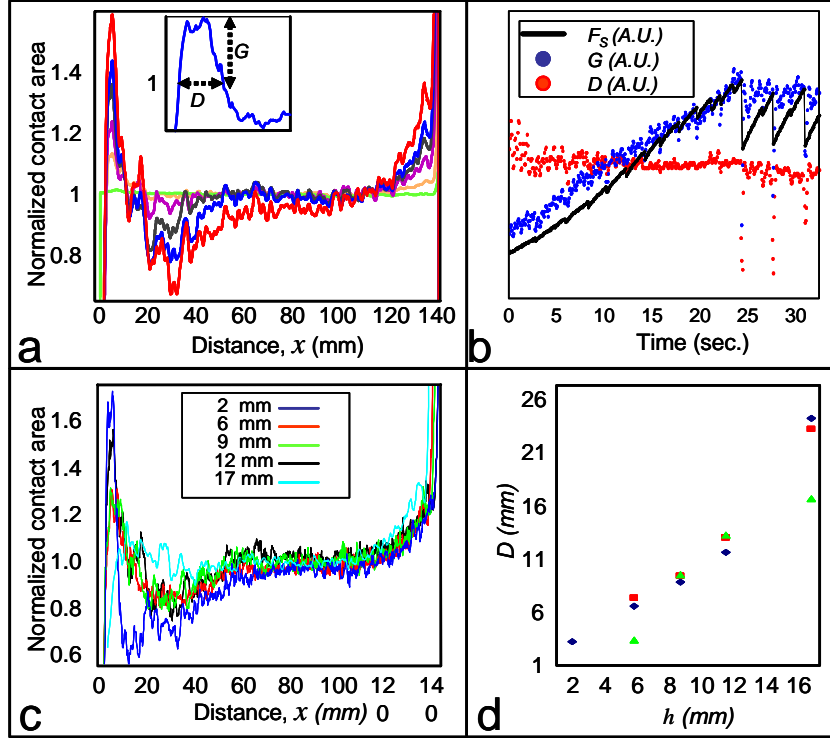


Figure 6: Evolution of the true contact area profile. **a**, $A(x,t)$ profiles formed by slip events of increasing length during a single experiment. Deformations of the initially uniform contact profile are both amplified and extended with each slip. (Inset) Close-up of $A(x,t)$ profile adjacent to the trailing edge depicting the width D and strength G of the asperity that is spontaneously formed there. **b**, F_S , D and G as a function of time for the experiment shown in (a) F_S , D and G are in arbitrary units to facilitate comparison. The asperity width D , stays nearly constant while G increases monotonically with F_S . **c**, D is determined h ; $A(x,t)$ in experiments where h was varied. The profiles are qualitatively similar in appearance. **d**, D , increases nearly linearly with h . All profiles in (c) were obtained for F_S at 75% of the value needed for the onset of stick-slip. $A(x,t)$ in **a** and **c**, is normalized with respect to its spatially uniform value, $A(x, t = 0)$, at the start of the experiment.

After overcoming this asperity, each slip both extends the length of the reduced contact area region created by its predecessors, and further reduces $A(x,t)$ by a significant amount. This systematic decrease of $A(x,t)$ creates a highly weakened region ahead of the asperity. As a result of this process, by the time large-scale overall motion (stick-slip) occurs, the established contact profile is highly non-uniform. As shown in Figure 6c, changes in h do not qualitatively affect

the general shape of the contact area profiles. Quantitative analysis reveals, however (Figure 6d), that the size of D is proportional to h .

One may suspect that the variations of the contact area is due to the torque imposed by the F_s , as F_s is applied at a finite height h rather than at $h=0$. We found that this effect is negligible over the range of used h . For example, the torque resulting from $h=2\text{mm}$ yields only a 3% variation of the normal stress over the interface length, whereas $A(x,t)$ varies by over 50% (Figure 6).

Surprisingly, once a contact profile is created, it is retained by the system, remaining nearly unchanged both after large-scale slip and in successive stick-slip events (RUBINSTEIN et al., 2007). Large internal stresses are, therefore, also retained by the system, even after major slip events occur. As the contact area mirrors the normal stress values, the existence of this non-uniform profile also indicates that the normal stresses along the interface (or fault) are highly non-uniform. In particular, the normal stresses along significant regions of the interface are considerably weaker than the remotely applied values would imply.

3.3. Simulations

We start our simulations in the state where the slider is unstressed, all surface contacts are intact and the entire interface is pinned. When the force is applied to the trailing edge of the slider, the shear stress accumulates in a finite region near the loading point. The solution of 1D elastic equations shows that displacements of the blocks, x_j , from their equilibrium positions, x_j^0 , and the corresponding contractions of the springs connecting the blocks, $x_j - x_{j+1}$, decrease exponentially with the distance from the trailing edge, x_0 , namely $x_j - x_j^0 \propto \exp\left[-\sqrt{K_s / NK} j\right]$. Thus, the length of the stressed region is determined by the ratio of the contact and slider stiffnesses, K_s / K . The exponential distribution of the shear stress along the slider is characteristic for the 1D model employed here, while a 3D description of the slider leads to a power law decrease of the stress (LANDAU and LIFSHITZ, 1986). Nevertheless, the localization of the shear stress in a finite range at the trailing edge that results from the 1D description is consistent with our experimental configuration where the finite spatial extent of the stressed region is determined by a height at which the force is applied.

As the applied force is increased, the stress in this region grows until it exceeds the thresholds for the rupture of surface contacts, f_{si} , and a detachment front starts to propagate across the interface (see Figure 7). The manifestation of the detachment fronts is seen in the loading curves, $F(t)$, presented in Figure 7a that reveals a sequence of small drops in $F(t)$. As in the experiments (e.g. Figure 3), these force drops correspond to discrete crack-like precursors to sliding

which occur well below the onset of overall motion and result from a minute motion of blocks at the slider trailing edge (see Figure 8).

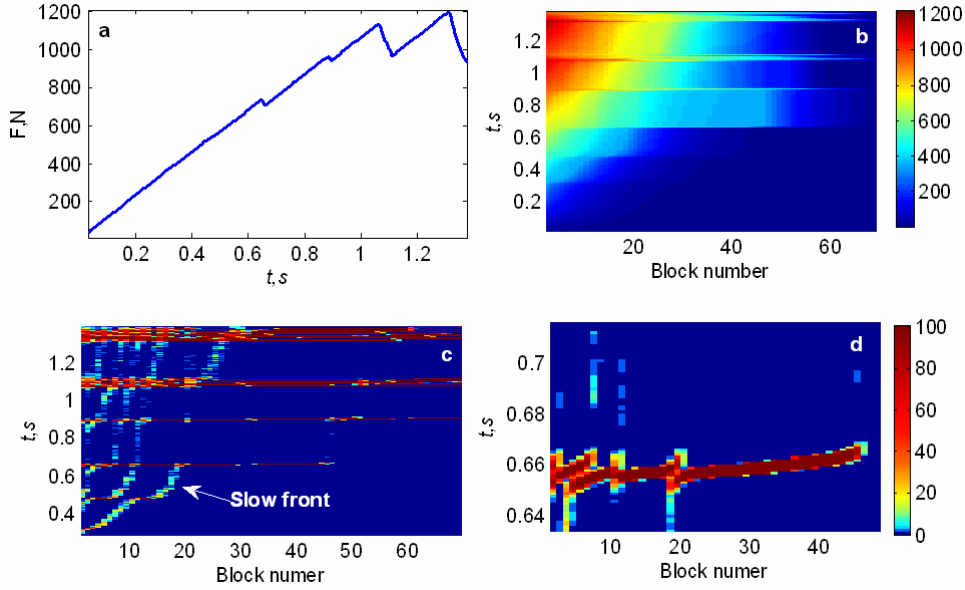


Figure 7: Onset of sliding for soft surface contacts, $K_s=K$. Panel (a) shows a loading curve, $F(t)$. Panels (b) and (c) present color maps for the distribution of elastic forces, $K(N-1)(x_j - x_{j-1})$ in the slider and the percent of detached contacts as functions of the block number, j , and time, t . Panel (d) presents an enlarged view of the fast detachment front from the panel (c) showing an excitation of a secondary sub-Rayleigh front by the slow fronts. The unstressed and stressed regions (b) and the regions with attached and detached contacts (c,d) are displayed by blue and red colors, respectively. The bars to the right of the maps set up a correspondence between the colors and the values of the force in Newtons (b), and the fraction of detached contacts in % (c).

What is the nature of the detachment fronts and their effect on the state of the system? In Figure 7 we present both the stress distribution along the chain of blocks (Figure 7b) and the fraction of detached contacts, as functions of the block number and time (Figure 7c). As in the experimental observations, we found that the onset of sliding is preceded by well-defined detachment fronts which are initiated at the trailing edge and extend over limited lengths across the slider that are smaller than its overall length. Figure 7b shows that these fronts generate a strongly nonuniform stress distribution across the interface. As a result, a new detachment front will propagate into an already highly stressed region that has been prepared by the previous front. The new front easily ruptures the pre-stressed contacts in this region, further extends itself and causes further elastic deformation of the slider. The threshold values of the applied force corresponding to the detachment fronts are considerably lower than the value needed to initiate overall motion of the slider, because only regions of limited length are fractured during these precursor events.

As in the experimental observations, our simulations show that three different types of detachment fronts play a role in the onset of sliding: (i) sub-Rayleigh (limited by the Rayleigh wave speed) fronts, (ii) slow detachment fronts, and (iii) super-shear (or “inter-sonic”) fronts that propagate at speeds beyond the shear wave speed. As a result of non-uniform loading, the precursors to sliding are always initiated at the trailing edge by the sub-Rayleigh front that rapidly accelerates until approaching the sound velocity, $V \propto \sqrt{K/M}$. This front is characterized by the simultaneous motion of a number of blocks which are detached from the base. Therefore, the velocity of the front, V , is determined by the elasticity of the slider and is independent of both the stiffness and rupture thresholds of the surface contacts. The properties of the contacts influence the number of blocks (of the slider domain) involved in a simultaneous motion, and the local displacements of the blocks. These decrease with the ratio between the stiffnesses of the surface contacts and the slider, K_s / K (see Figure 7). Figure 7 also shows that, during transition to sliding, the blocks perform stick-slip motion where slips correspond to the precursor events, and the slip lengths are in the micrometer range.

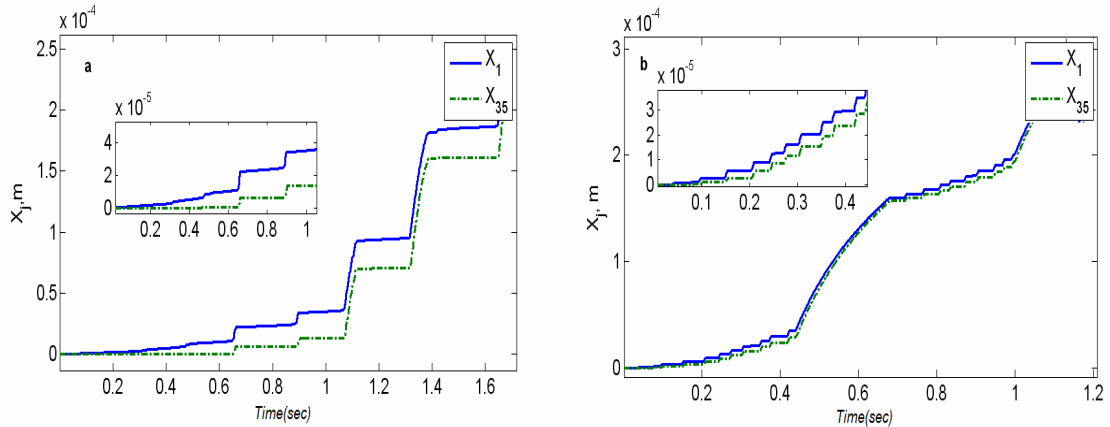


Figure 8: (color on line) Displacements of the first block, x_1 (trailing edge) and the middle block, x_{35} , as a function of time for the case of soft (a) and stiff (b) surface contacts. Insets present enlarged views of the motion of blocks prior the onset of overall sliding.

Once the sub-Rayleigh front extends beyond the highly-stressed region, it arrests. At the tip of the arrested front, however, a high excess stress is retained. This stress focusing is expected at the tip of an arrested rupture. The accumulation of stress is clearly seen in Figure 9, where we present a 2D map for the stress distribution in the surface contacts as functions of the block number and time. Here we show that the average force experienced by contacts connecting a given block j to the base, $f_j = \frac{1}{N_s} \sum_i^{N_s} f_i^{(j)}$, where $f_i^{(j)} = k_i^{(j)} l_i^{(j)}$, $l_i^{(j)}(t)$ is the contact elongation, and the averaging is taken over all contacts corresponding to the same block j . This excess stress triggers a slow detachment front that is similar to those observed experimentally (e.g.

Figure 4a-bottom). As in the experiments, these slow fronts propagate with a velocity that is over an order of magnitude lower than V_R (see Figures 7c and 9). In contrast to the sub-Rayleigh fronts, which exhibit the collective motion of the detached blocks, the slow fronts represent the motion of individual blocks. At each moment only one or a few blocks become partially detached from the base and move. As a result, the velocity of the slow front is mostly determined by the frictional properties of the individual blocks (stiffness of surfaces contacts, thresholds of rupture forces) and only slightly depends on the slider elasticity. In particular, the velocity decreases with the increase in interfacial stiffness, K_s , and mean rupture force, F_s .

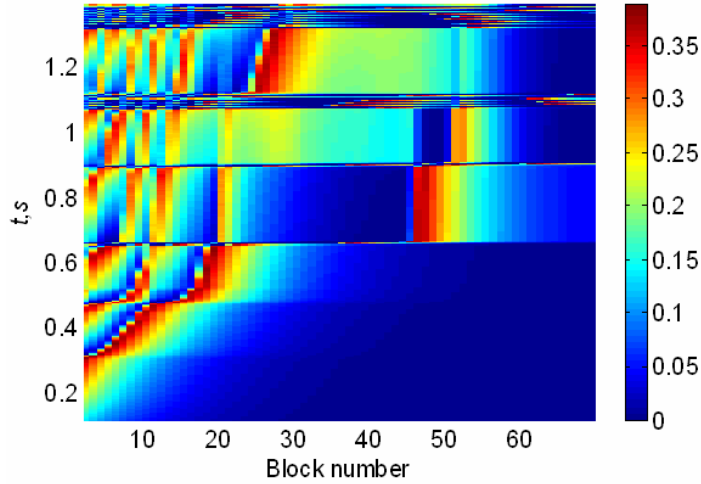


Figure 9: Color map for the distribution of elastic forces, $f_j = \frac{1}{N_s} \sum_i^{N_s} f_i^{(j)}$, in the contacts as functions of the block number, j , and time, t . The unstressed and stressed regions are displayed by blue and red colors, respectively. As in Figure 7, these simulations were performed for $K_s \sim K$. The bar to the right of the maps set up a correspondence between the colors and the values of the force in Newtons.

Figures 7c and 7d show that the interaction between the sub-Rayleigh front propagating from the trailing edge and the slow fronts triggers the excitation of secondary sub-Rayleigh fronts at various distances from the edge. This effect leads to a formation of a new type of fast fronts which corresponds to a superposition of sub-Rayleigh fronts that are initiated simultaneously at different locations across the slider. An effective velocity of such fast fronts can be few times higher than V_R . These “chained” fronts may be analogous to the super-shear (inter-sonic) detachment fronts that have been observed experimentally (RUBINSTEIN et al., 2004, 2007), under the loading conditions noted in Figures 1 and 2, as well as under either impact loading conditions from the system’s edge (ROSAKIS et al., 1999, 2000) and loading conditions for which fronts were triggered explosively (XIA et al., 2004). This overall scenario for super-shear rupture may be rather general, as finite element simulations (NEEDLEMAN, 1999) performed for the

case of shear fracture under impact loading (ROSAKIS et al., 1999, 2000) suggest a similar scenario, in which the transition to super-shear rupture involves micro-crack nucleation ahead of the main rupture front.

Figure 7b demonstrates that with each successive detachment front, the stress distribution across the slider becomes increasingly more nonuniform. By the time the system is ready to slide, the precursor fronts have generated highly stressed state of the slider and the contacts. Hence, the transition to sliding occurs in the highly nonuniform, stressed system. The transition to sliding is manifested by a significant drop in the loading curve, $F(t)$, that is an order of magnitude larger than the small drops corresponding to precursors of sliding. Contrary to the precursors which are caused by the discrete detachment fronts, the transition to overall sliding occurs through a quasi-continuous (in time) set of fronts propagating one after another from the trailing to the leading edges of the slider. During the time interval corresponding to this transition, most surface contact are simultaneously detached and, therefore, allow the overall motion of the slider.

It is notable that for the system with moderately flexible surface contacts ($K_s \approx K$) the nonuniform stress distribution produced by the sequence of detachment fronts prior to the first sliding event remains virtually unchanged in the subsequent stick-slip motion. This result is consistent with experimental observations (RUBINSTEIN et al., 2007) which suggested that nonuniformity of the contact is the preferred state of the system during sliding.

Our calculations demonstrate that dynamics of the transition to sliding depends on the material properties. For instance, for stiff surface contacts with $K_s = 50K$ the slow fronts do not evolve, since in this case the excess stress accumulated at the tip of the arrested sub-Rayleigh front cannot overcome the resistance of the surface contacts and initiate slow motion (see Figure 10). As a result, in this case we also did not find the super-shear fronts (see Figures 10b and 10d), and the onset of sliding is completely determined by the propagation of the sub-Rayleigh fronts. Another distinctive property of a surface with stiff contacts is a release of high stresses in the slider and renewal to nearly a uniform stress state that is precipitated by sliding. The difference in stress relaxation during the stick-slip transition leads to a very different mechanism of stick-slip motion for soft and stiff surface contacts. In the case of soft contacts the second and subsequent stick-slip events occur at the pre-stressed interface and do not involve a sequence of crack-like precursors (see Figures 7 and 8). This scenario is similar to the experimental observations.

In the case of stiff contacts where the stress is relieved by sliding, all stick-slip transitions occur through the excitation of discrete detachment fronts (see Figures 8b and 10). Thus dynam-

ics of transition from static to kinetic friction strongly depends on the ratio between the stiffnesses of the surface contacts and the slider, K_s/K .

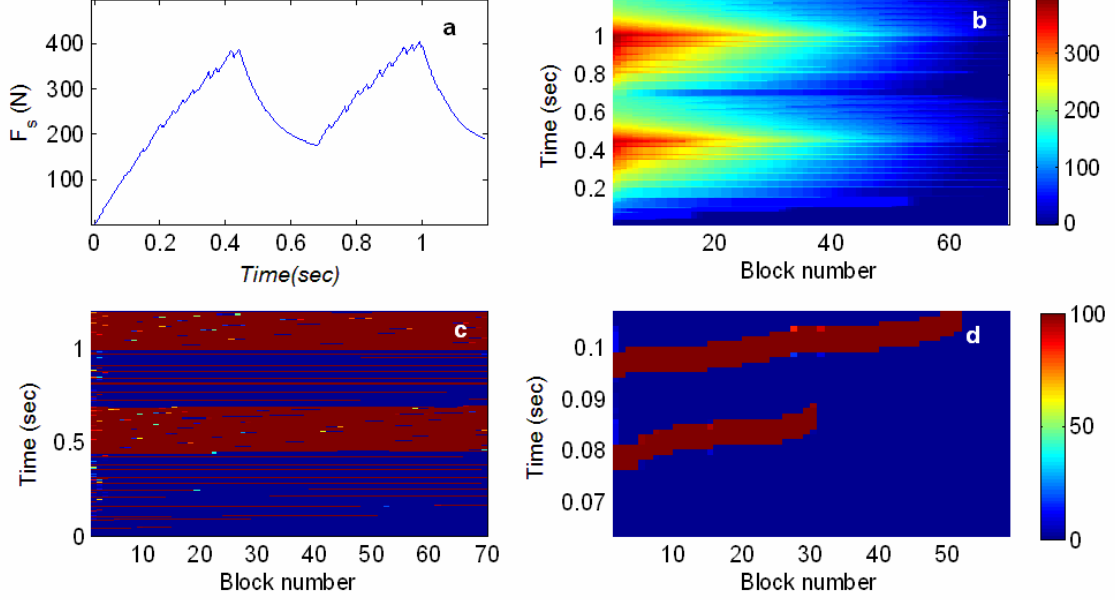


Figure 10: Onset of sliding for stiff surface contacts, $K_s=50K$. Panel (a) shows a loading curve, $F(t)$. Panels (b) and (c) present color maps for the distribution of elastic forces, $K(N-1)(x_j - x_{j-1})$ in the slider and the fraction of attached contacts as functions of the block number, j , and time, t . Panel (d) presents an enlarged view of the detachment front from the panel (c) showing a propagation of a sub-Rayleigh front. The notation is as in Figure 7.

3.4 Synthesis of experimental and theoretical results

These experiments and simulations suggest an intuitive picture for the sequence of events leading to frictional sliding. Before the onset of slip, the loading at the trailing edge imposes a high shear stress region near the edge, whose magnitude decays over a length proportional to h . When F_s is sufficiently large, this highly stressed region yields and an initial crack-like slip event is generated. The slip traverses this region and arrests at a distance, l_0 , where the shear stress level is below the slipping threshold. This event results in: (1) slip within the region l_0 , where built-up shear stresses are released, (2) elastic deformation of the slider to compensate for the slip-induced contraction in the x direction. The deformation results in an outwardly protruding region of size $D \sim 1/3l_0$ in which the contact area increases, thereby dynamically forming an asperity (cf. Figure 6a). The inwardly bowed region over the remainder of the region l_0 , reduces the normal force (equivalent to a decrease of $A(x,t)$) (3) establishment of a high residual shear stress concentration entrained in the vicinity of the point of arrest of the slip event (This is due

to the stress singularity that occurs at the tip of a shear crack.). Upon further increase of F_s , the barrier imposed by the asperity is again overcome and a new slip event is generated. This slip event will easily traverse the weakened region beyond the asperity. Arriving at the tip of the previously arrested front, the new slip event will add sufficient energy to release the energy stored within the high-stress region imposed previously, thereby enabling it to extend itself by Δl . This extension is accompanied by further elastic deformation of the block, thereby increasing the non-uniformity of $A(x,t)$. In this way, each slip event transfers the shear stress imposed at the boundary further along the interface.

4. DISCUSSION

The influence of fault complexity on the dynamics of rupture propagation and seismicity has been the subject of intensive past and recent research (AKI, 1979; DAS, 2003; LAY et al., 1982). Our experiments suggest that the fingerprint of an ‘edge’ can be traced, not only to the dynamics of a given rupture, but throughout the seismic cycle of a given fault. The results suggest that a geometrical inhomogeneity introduces a scale that may govern the nucleation, size and repeatability of earthquakes along a given fault.

The experiments and model presented here paint an interesting picture of earthquake dynamics along faults that are loaded at an edge or a step. They indicate that stress transfer along such faults may be mediated by a periodic sequence of precursory events. This sequence of slip events culminates by the triggering of slowly propagating front, which leads to system-size events. The experiments suggest that the early slip events of the periodic sequence (within the scaling regime shown in Figure 5a) “feel” an effectively infinitely long (unsegmented) fault. In contrast, the accelerated growth of l that marks the breakdown of scaling prior to the onset of large events indicates that the dynamics are affected by the fault size during the nucleation phase (OHNAKA and SHEN, 1999) of a large event. This accelerated increase in l is strongly suggestive of the accelerated seismic release that precedes some large earthquakes (BUFE and VARNES, 1993). The results also imply that precursory sequences of events that initiate from a fault edge strongly modify the fault contact plane prior to a large event.

One example of such slip sequences along a fault edge in the crust is portrayed by the foreshock sequence of the 1998 Sendai Bay event, along the Nagamachi-Rifu fault, Japan (UMINO et al., 2002). In this field case, the main shock of M 5.0 was preceded by 17 foreshocks ranging in magnitude from 1.7 to 3.8 with essentially identical seismic characteristics. The foreshocks sequence lasted about three days with the largest foreshock occurring six minutes prior to the main shock. Mechanical modelling of this sequence of events suggests that the fault was edge-loaded by non-seismic slip in the lower crust (NAKAJIMA et al., 2006). Accordingly, the fore-

shock hypocenters propagated upward along the locked part of the Nagamachi-Rifu fault. Finally, Umino et al noted that “A small ambiguous phase...is observed in seismograms of both the M5.0 main shock and the M3.8 largest foreshock...” (UMINO et al., 2002). This slow, low amplitude ambiguous phase is lacking in the other foreshocks and all aftershocks, and is likely the equivalent of the slow fronts observed in our experiments immediately *before* the main slip event (Figures 4a and 7). These slow fronts may also relate to the accelerated creep events that are anticipated to be part of the nucleation phase of major earthquakes (DIETERICH and KILGORE, 1996; OHNAKA and SHEN, 1999). Thus, in spite of scale and complexity differences, we note the following similarities between the 1998 Sendai Bay events and our experiments and simulations: (1) qualitatively similar (edge) loading conditions, (2) a distinct sequence of precursory events; (3) initiation of precursory events from nearly the same location and (4) a slow (“ambiguous”) phase that occurs only before the main event.

Another outstanding example of a sequence of foreshocks is the sequence of events that preceded the 1995 Mw 7.9 off-Etorofu earthquake, Kurile Islands (HURUKAWA, 1998). We show below that our experimental results can serve as laboratory analog for the foreshock-mainshock sequence of this earthquake. Hurokawa (HURUKAWA, 1998) relocated the foreshocks, the mainshock, and aftershocks, and used quality criteria (e.g., number of recording stations, minimum magnitude and distance, and low standard error), to analyze a total of 238 high-quality events. These events occurred along a 14°-19° dipping thrust of the boundary between the North American and Pacific plates (Fig. 11A). A series of 51 foreshocks initiated nine days before the mainshock, including three events of $M > 6$. The foreshock area extended from depths greater than 40 km in the direction of the trench axis (Fig. 11B), with the final area covered by foreshock activity roughly 80×30 km in size (Fig 11C). The mainshock rupture initiated at the deepest point of the foreshock area, and its final area, as determined from aftershock distribution, was roughly six times the final area of the foreshock distribution (Fig. 11C).

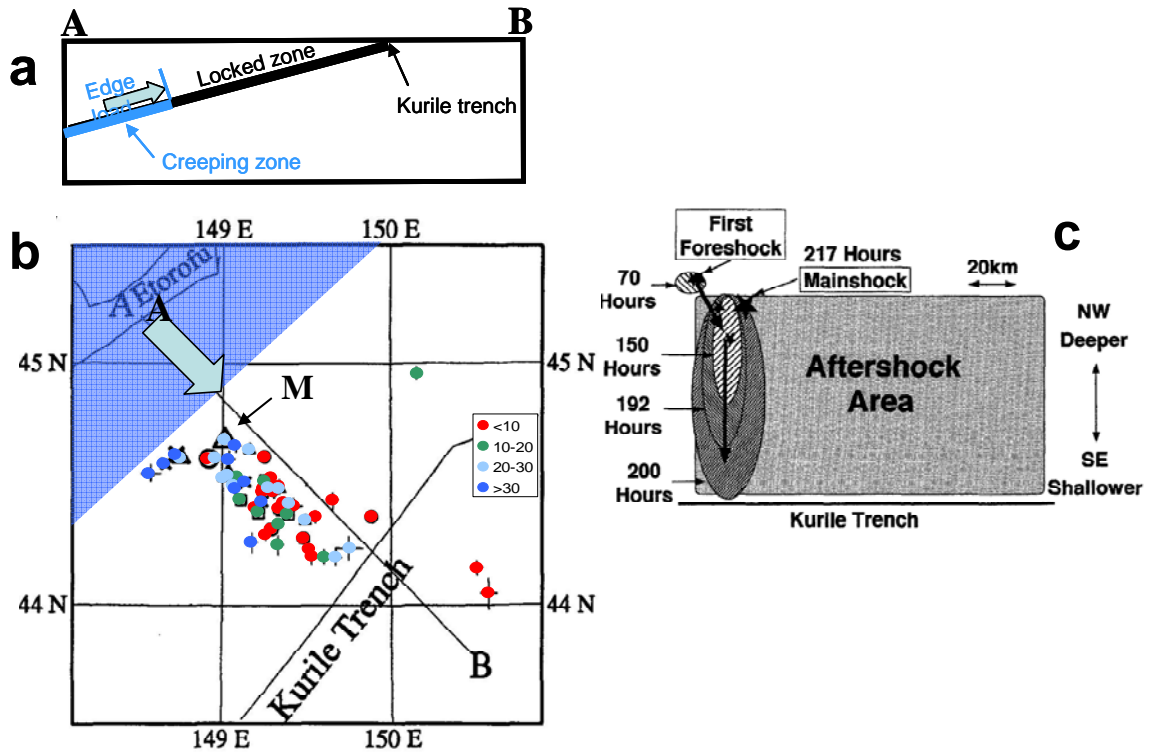


Figure 11. The foreshock sequence of the 1995, M7.9 off-Etorofu earthquake, Kurile trench; data from (HURUKAWA, 1998). (a) A schematic 2D model of the Kurile subduction zone in the area of the earthquake; the fault dips 12° - 19° to the NW (HURUKAWA, 1998). We suggest that the upper, locked part of the fault (black line) is edge-loaded by the creeping lower part (blue). (b) Map of relocated foreshocks and mainshock (modified Fig. 5B in (HURUKAWA, 1998)). Events are divided into four depth groups marked in color (legend in km). The area in blue in the NW corner is the creeping zone according to the model in A; M denotes the mainshock. (c). Figure 8 from (HURUKAWA, 1998) showing a schematic distribution of the foreshock activity (see text). The arrows show the dominant propagation direction. Note that the mainshock hypocenter is at the deepest edge of the foreshock area; stars represent magnitude 6 or larger earthquakes.

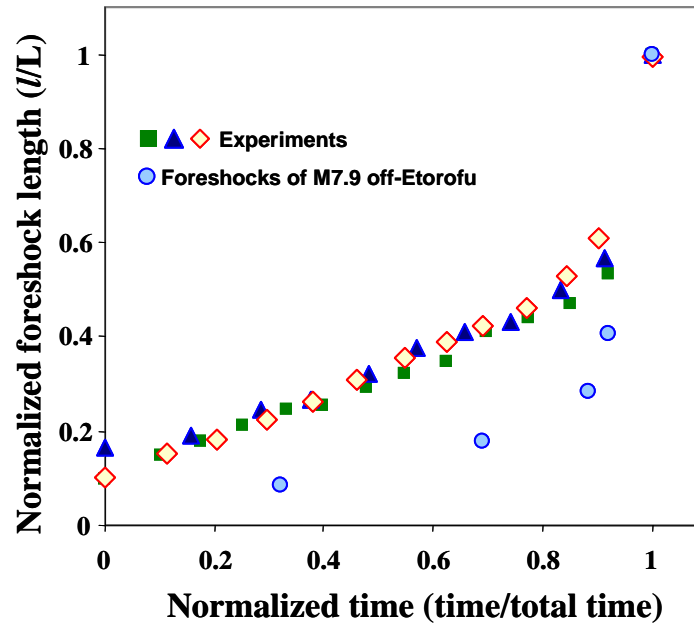


Figure 12. Time-progression of arrested slip events in three typical experiments (square, triangle, diamond symbols) of the present study and in the foreshock sequence (circles) of the 1995, M7.9 off-Etorofu earthquake, Kurile trench. The latter data were obtained from Figure 11c (Figure 8 in (HURUKAWA, 1998)) and modified, as explained in the text. The linear temporal dependence in the experimental data reflects the constant forcing rate used in the experiments. The forcing rate in the Etorofu events is the (unknown) creep rate at depths over 50km. This may be the reason that the field data exhibit nonlinear temporal dependence of the foreshock sequence.

We believe that the above observations of HURUKAWA (HURUKAWA, 1998) are analogous to the present experiments. This statement is based on the following similarities:

Edge-loading. In the experiments, the laboratory “fault” was loaded by a continuously increasing force, F_s , at its edge (c.f. Figure 1). In the field, the upper, brittle, locked part of the Kurile Islands thrust at shallow depths < 50 km (black, inclined line in Figure 11A), was loaded by its deeper, ductile, creeping part at depths > 50 km (blue inclined line in Figure 11A and blue area in Figure 11B).

Space-time progression of slip events. In each experimental run, a series of arrested events propagated from the loading edge and extended the length of the slipped region (Figure 3). This series of arrested events was culminated by an un-arrested event along the entire block (Figure 3). We regard this event as the analog of a mainshock. In the off-Etorofu events, the foreshocks initiated in the deepest part of the brittle fault and propagated southeastward (above and Figure 11B). As in the experiments, the mainshock also initiated at the loaded edge (the deepest part), covered the foreshock area and extended laterally (Figure 11C).

Increase of the affected length (area). Both the experiments and the field display systematic temporal increase of the affected length (along experimental fault, c.f. Figures 3 and 5) or af-

affected area (Kurile Islands thrust, Fig. 11C). This temporal progression is presented in Figure 12 for the normalized values of length (area) and time. For the experiments, the normalized affected length is the ratio between the length of the arrested event, (l in Figure 4) and the length of the un-arrested event (L is length of the experimental fault shown in Figure 1). For the field data, the normalized affected length is $\sqrt{A_F / A_M}$ where A_F is the area affected by the foreshocks at a given time, and A_M is the area of the mainshock; both A_F and A_M were determined from Figure 11C, which is the model of Hurukawa (c.f. Figure 8 in (HURUKAWA, 1998)). We use the square root of the area ratio to facilitate a comparison between area in the field and length in the experiments. The normalized time is $t_n = (t - t_0) / \Delta t$, where t is the time of the event (experimental or foreshock), t_0 is the time of the first event, and Δt is the time difference between the first event and the last. For the experiments, we use the results of three runs loaded by similar rates but under different normal stresses.

Figure 12 displays general similarities between the time-progression of the arrested events in the experiments and the foreshocks of the 1995 Mw 7.9 off-Etorofu earthquake. In both settings, the length, l , (area) systematically increases with time for almost the entire duration of the sequence, until reaching approximately 40%-45% of the final length, L (area). L is analogous to the additional length (area) is covered during the mainshock in the field. In the figure, the linear progression of l in the experiments is due to the linear increase with time of the applied shear force, F_s . In the off-Etorofu sequence, we don't know whether the creep (edge-loading) rate was constant over the duration of the foreshock sequence. If so, we would expect the same quasi-linear temporal increase of the event sizes.

We speculate that the general similarities between the experimental sequences of slip event and the seismic foreshock sequences of off- Etorofu and Sendai earthquakes are not coincidental. We believe that this similarity reflects the effect of edge-loading on a locked fault, and event progression that is needed to weaken the locked segment and to prepare it for the mainshock. As both our model and experiments show, the qualitative (generic) form imbued in edge-loading is *sufficient* to trigger this general sequence of events. We, therefore, expect that a locked fault segment which is edge-loaded by a deep, creeping segment, e.g. in subduction zones, is likely to fail in a mode similar to our experiments.

In conclusion, the results of this study have shown that the fact that shear is applied non-uniformly to a sliding system leads to complex, *systematic* behaviour that appears analogous to natural phenomena, whose source is currently not well understood. We believe that the analogies between our experimental results and seismic observations stem from their similar edge-loading configurations. As edges and bumps are common in faults at many scales (SAGY et al., 2007; STEIN et al., 1997), it is therefore anticipated that this loading will generate stress distri-

butions that are similar to the laboratory model, and, consequently, may lead to similar dynamics.

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